

Supporting Information for "Uncovering the geodetic signature of silent slip through repeating earthquakes"

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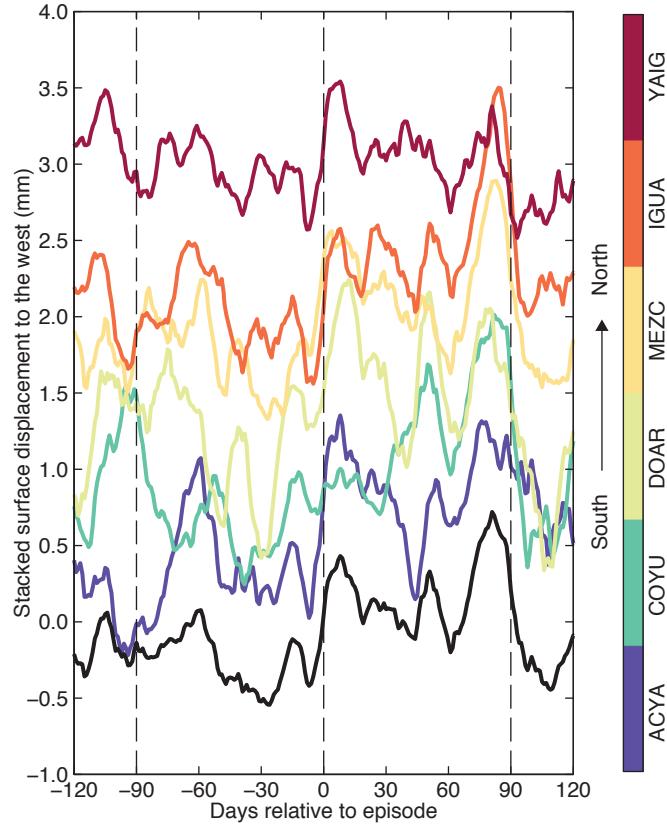


Figure S1. Stacked smoothed west-east displacement time series on the regional cGPS network. The average stacked displacement for the plotted subset of stations is shown in black. For a more complete description, please see Figure 3 of the main text.

Text S1.: Static inversion for the slip distribution at the plate interface The displacement vectors are determined by applying a 10-day moving average to the stacked GPS time series and then calculating the difference between ± 10 days. The GPS surface displacements are then inverted to evaluate the slip distribution on the subduction interface. The modeled fault plane that represents a simplified subduction interface extends 468 km along strike. The plane has a dip angle of 15° and becomes subhorizontal 150 km inland, with a depth of 40 km [Radiguet et al., 2012]. The Green functions were calculated for a layered elastic half-space as-

suming a layered crustal model [Hernandez et al., 2001]. The static GPS displacements are inverted using a linear problem least-square formulation [Radiguet et al., 2011, 2012] with an additional positivity constraint. The rake of the inversion is fixed, and optimized for lower misfit, resulting in a rake angle of 90° (pure thrust) for the 2006 SSE inversion and 100° (in the direction of the plate convergence) for the small slow slip. Smoothing is controlled by a correlation length, describing the correlation between nearby fault patch fixed to 50 km for the 2006 SSE and to 200 km for small slow slip.

References

- Hernandez, B., N. M. Shapiro, S. K. Singh, J. F. Pacheco, F. Cotton, M. Campillo, A. Iglesias, V. M. Cruz-Atienza, J. M. Gomez, and L. Alcantara (2001), Rupture History of September 1999 Intraplate Earthquake of Oaxaca, Mexico ($M_w=7.5$) from Inversion of Strong-Motion Data, *Geophys. Res. Lett.*, 28(2), 363–366, doi:10.1029/2000GL011975.
- Radiguet, M., F. Cotton, M. Campillo, B. Valette, V. Kostoglodov, and N. Cotte (2011), Spatial and temporal evolution of a long term slow slip event: the 2006 Guerrero Slow Slip Event, *Geophys. J. Int.*, 184(2), 816–828, doi:10.1111/j.1365-246X.2010.04866.x.
- Radiguet, M., F. Cotton, M. Vergnolle, M. Campillo, A. Walpersdorf, N. Cotte, and V. Kostoglodov (2012), Slow slip events and strain accumulation in the Guerrero gap, Mexico, *J. Geophys. Res.*, 117, B04305, doi:10.1029/2011JB008801.

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